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ON THE SPATIAL RELATIONSHIP OF 1-METER EQUATORIAL IRREGULARITIES AND DEPLETIONS IN TOTAL ELECTRON CONTENTS

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1 February 1979

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20. ABSTRACT (Continue on reverse side if necessary and identify by block number) A radar experiment was conducted at Kwajalein Atoll, Marshall Islands to investigate the spatial relationship of 1-m equatorial spread-F irregularities to total electron content (TEC) depletions. A high-power radar was operated (1) in a backscatter scan mode to spatially map the distribution of 1-m irregularities, and (2) in a dual-frequency, satellite-track mode to obtain the longitudinal TEC variations. Using the radar data, we show that radar backscatter "plumes" found in the disturbed, nighttime equatorial		

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20. ABSTRACT (Continued)

Ionosphere are longitudinally coincident with TEC depletions. We suggest that the TEC depletions are probably due to the presence of plasma bubbles in the equatorial F layer.

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PREFACE

The author would like to thank Dr. D. M. Towle, MIT Lincoln Laboratory, for making the ALTAIR data available and for continued co-operation during the project.

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I INTRODUCTION

Recent research on spread-F phenomena in the nighttime equatorial ionosphere has led to the discovery of two unique features, plasma bubbles and radar backscatter plumes. Plasma bubbles are "biteouts" (up to 200 km in diameter) in the F-region ionosphere with electron density depletions up to three orders of magnitude below the ambient level. Radar plumes are regions of strong backscatter produced by field-aligned irregularities (FAI) that extend upward from the bottomside to the topside of the F layer. Plasma bubbles were first observed in ion concentration measurements made by the OGO-6 satellite (Hanson and Sanatani, 1973). Since then, biteouts in plasma density have been detected in-situ by rockets (Kelley et al., 1976; Morse et al., 1977) and by other satellites (McClure et al., 1977; Dyson and Benson, 1978). Radar backscatter plumes have been mapped at 50-MHz by the Jicamarca radar (Farley et al., 1970; Woodman and La Hoz, 1976; Basu et al., 1977), and at 155 MHz by the ALTAIR radar (Tsunoda et al., 1979).

Plasma bubbles have been explained in terms of the Rayleigh-Taylor instability (Dungey, 1956; Hudson and Kennel, 1975). Numerical simulations of the nonlinear evolution of the collisional Rayleigh-Taylor instability by Scannapieco and Ossakow (1976) have shown that plasma depletions develop in the bottomside of the F layer and rise into the topside of the F layer. Furthermore, because of a dominant nonlinearity in two-dimensional models, the bubbles become highly elongated in the vertical direction (Chaturvedi and Ossakow, 1977; Hudson, 1978).

Because radar backscatter plumes are similar in size, shape, and behavior to plasma bubbles, researchers have proposed that the two phenomena are directly associated, if not spatially coincident (Woodman and La Hoz, 1976; McClure et al., 1977). However, thus far, there has been no experimental evidence verifying this hypothesis.

In this report, we present the findings of a radar experiment in which the longitudinal variations of total electron content (TEC) were mapped in the vicinity of radar backscatter plumes. We show that radar backscatter plumes are longitudinally coincident with TEC depletions found in the nighttime equatorial ionosphere. On the basis of evidence presented by other workers, we suggest that the TEC depletions are probably caused by the presence of plasma bubbles.

II THE EXPERIMENT

Both radar backscatter plumes and TEC variations were spatially mapped using ALTAIR, a high-power radar, located in the Kwajalein Atoll, Marshall Islands. A description of ALTAIR and its capabilities for equatorial spread-F studies were published by Tsunoda et al. (1979). In contrast to the Jicamarca radar which uses a more or less fixed antenna beam, ALTAIR utilizes a fully-steerable 46-m paraboloid antenna to spatially map equatorial field-aligned irregularities (FAI) in the east-west and near-vertical directions. The angular sector scanned by ALTAIR was 45° in the geomagnetic east-west direction, or approximately 300 km in lateral extent at 400 km altitude.

In addition to mapping equatorial spread-F irregularities at 155.5 MHz, ALTAIR is capable of "skin-tracking" satellites of opportunity at two frequencies (155.5 and 415 MHz) simultaneously. The difference in range (group delay) to the satellite, measured at the two frequencies, provides an absolute measure of the total electron content. The longitudinal distribution of TEC can be determined by tracking east-west orbiting satellites.

On 26 August 1977, ALTAIR was operated in a west-to-east sector-scan mode from 0930 to 1250 UT. (Local time in Kwajalein lags Universal time by 12 hours.) During most of the night a spatial map of backscatter strength was made every six minutes. These scans were interrupted between 1032 and 1040 UT to skin-track the PEOLE (French) satellite. The satellite pass geometry relative to the ALTAIR scans is shown in Figure 1. The location of the point directly beneath the satellite, as a function of time, is shown by the line labeled sub-satellite track. The sub-satellite track is seen to be in a generally west-to-east direction, with an orientation that is slightly north of the east-west plane. The elevation angle to the satellite at the time of closest approach to ALTAIR (1034:54 UT) was 76.7° .

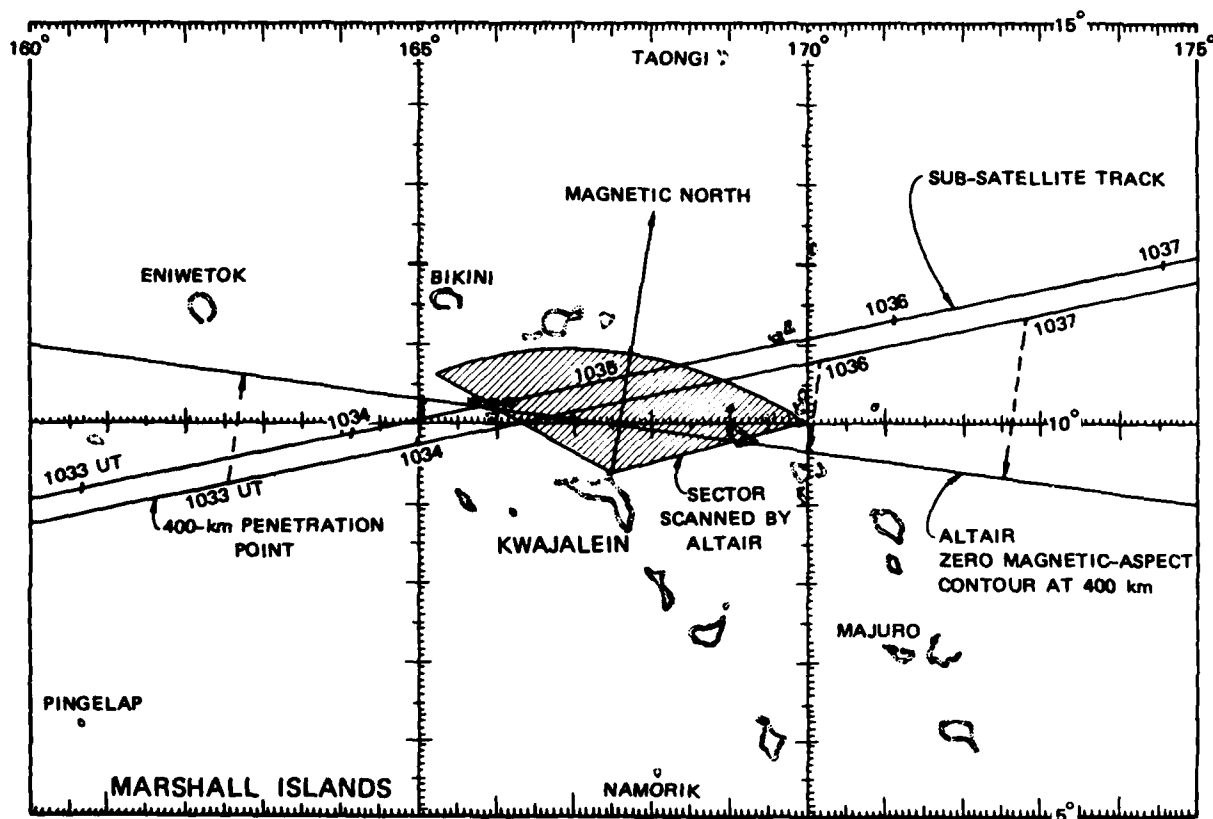


FIGURE 1 SATELLITE PASS GEOMETRY RELATIVE TO THE ALTAIR SCANS

Because the principal ALTAIR radar plumes (described below) to which we compared the TEC variations were centered at an altitude of 400 km, we converted the temporal TEC variations to spatial ones by using that altitude as a reference. The 400-km penetration point (from satellite to ALTAIR) as a function of time is shown in Figure 1 by a line parallel to the sub-satellite track.

Finally, in order to make a direct comparison of spatial TEC variations to backscatter plumes, we projected the TEC spatial variations along the 400-km penetration reference line onto the plane scanned by ALTAIR by assuming that both radar plumes and TEC variations are magnetic field-aligned. For example, the TEC measurements made at 1033 and 1037 UT were projected as shown by the dashed arrows from the 400-km penetration curve onto the geomagnetic east-west line in Figure 1, which is a ground projection of the zero magnetic-aspect contour at 400 km for ALTAIR. With this projection of the TEC variations, the depletions in TEC can be directly compared with the radar plumes mapped by ALTAIR. We note that the projection of the principal TEC variations is no more than 50 km along geomagnetic field lines. Morse et al. (1977) found strong correlation of ionospheric features up to a few hundred kilometers along magnetic flux tubes.

An important feature of this particular experiment geometry is that the PEOPLE satellite was at a nominal altitude of 550 km. This means that the TEC variations are produced by electron density variations within the limited altitude range from 550 km down to about 250 km, the altitude of the bottom of the F layer. Therefore, the TEC variations must result directly from the altitude regime containing radar plumes, and are not obscured by electron density variations that might have occurred well above the radar plumes.

III RESULTS

The results are summarized in Figure 2. In the upper panel we have plotted the variation of TEC as a function of magnetic east distance from ALTAIR. (TEC data were available at 3-second intervals, or approximately every 20 km in ground distance.) The times of TEC measurement at given distances from ALTAIR are also shown (above the upper panel). The measured TEC variations are plotted as the irregular boundary of the shaded regions while the smooth boundary is the TEC variations expected from a spherically stratified ionosphere. The smooth curve is included to facilitate identification of TEC depletions and other deviations in the irregular curve. Although the smooth curve was selected arbitrarily, it is consistent with estimates of peak electron density in the quiet F layer (defined as regions with no radar backscatter) computed from ALTAIR incoherent-scatter (IS) measurements. That is, the vertical TEC value is 10^{17} el/m² for the smooth curve. If we assume that the applicable layer thickness for the vertical TEC measurement is 300 km, we obtain a mean electron density of 3×10^5 el/cm³. The ALTAIR IS measurements indicate a value of $3-4 \times 10^5$ el/cm³. The shaded regions therefore represent TEC depletions and enhancements presumably associated with plasma bubbles and a generally disturbed equatorial ionosphere.

Major TEC depletions of 45% and 35% were seen directly over ALTAIR around 1035 UT. Other smaller TEC depletions are also seen (in Figure 2) to be distributed across the 1700-km sector scanned with the satellite. In addition to the depletions, there are also TEC enhancements that are as much as 20% above the smooth TEC curve.

The ALTAIR backscatter results are presented in the lower panel of Figure 2. The spatial distribution of field-aligned backscatter was constructed as follows. The 1-m FAI regions were assumed to drift eastward at a nominal speed of 100 m/s (Woodman, 1972). The assumed drift speed was used to space the backscatter maps, taken over ALTAIR at

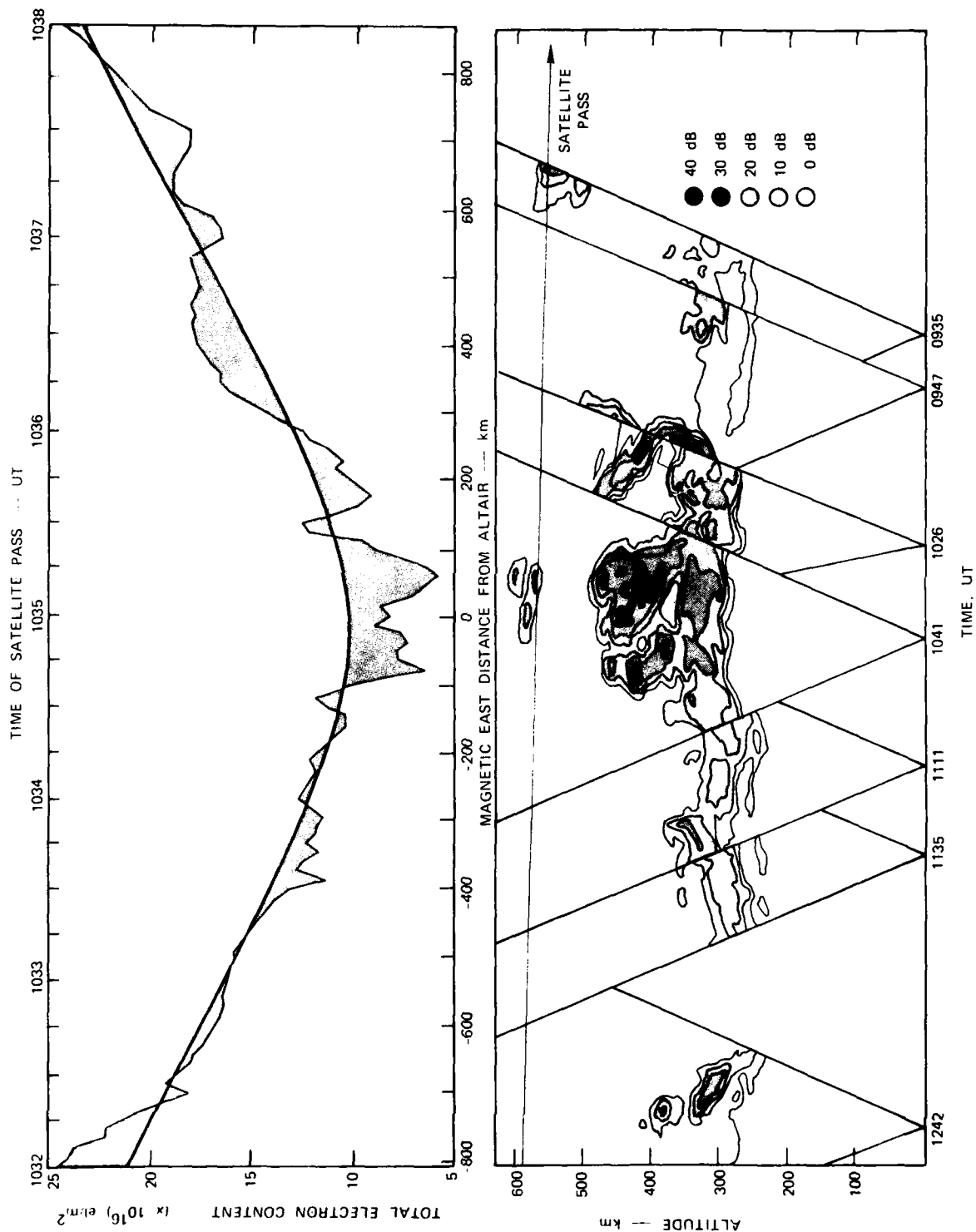


FIGURE 2 COMPARISON OF TOTAL ELECTRON CONTENT VARIATIONS AND ALTAIR BACKSCATTER PLUMES

different times, in the pattern shown in Figure 2. The position of each map, relative to the location of ALTAIR, was based on the time of closest approach of the PEOLE satellite. As can be seen from Figure 2, the TEC measurement over ALTAIR was made at 1035 UT and the corresponding backscatter map was made at 1040:40 UT. The westward displacement of the map is 30 km, based on the assumed 100 m/s eastward-drift speed.

In this analysis, we were primarily interested in comparing the TEC variations measured directly over ALTAIR to the corresponding spatial distribution of radar backscatter plumes. The reason for restricting the analysis to overhead data is that several complications arise when one attempts to make comparisons elsewhere. Comparisons of TEC features located further east and west of ALTAIR require increasing amounts of temporal extrapolation. (We present evidence later in this section that suggests the lifetime of radar backscatter regions is not more than several tens of minutes.) Furthermore, TEC measurements are integrated estimates of the electron density along the line of sight. As a result, low-elevation TEC measurements cannot be simply related to longitudinally-confined but vertically-extended radar plumes.

The spatial relationship between the radar plumes overhead of ALTAIR at 1035 UT and the TEC depletions, shown in Figure 2, is striking. The plumes are seen to be spatially coincident with the TEC biteout that extends from -100 to +100 km from ALTAIR. Even the recovery (at 1035 UT) within the TEC depletion, centered directly over ALTAIR, corresponds to a region of very weak backscatter. The second major TEC depletion (at 1035:45 UT) located at 200 km east of ALTAIR is seen to be also coincident with backscatter that extends upward in altitude (see map taken at 1026 UT). The spatial coincidence of backscatter plumes and TEC biteouts also appears to hold true in the features to the west of ALTAIR, although the relationship is not as convincing as in the overhead comparison. The features that appear to the extreme east and west distances from ALTAIR should not be critically compared because of the required temporal extrapolation and the low elevation angles.

It is also interesting to note that the enhancements in TEC appear to correspond to regions of very weak or of no backscatter. The case in

point is the TEC enhancement located between 300 and 500 km east of ALTAIR. The TEC enhancement at the extreme west distance is also supportive of this relationship.

The conclusion that backscatter plumes are longitudinally coincident with TEC depletions can be strengthened by noting that the patchy backscatter associated with the TEC depletion observed at 1035:45 UT was in fact an intense backscatter plume at earlier times. The time evolution of the backscatter plume is shown in Figure 3. We see that at 0947 UT, there existed an intense, backscatter plume that extended upward to an altitude of 600 km. The next map, taken 39 minutes later at 1026 UT, shows the dissipation of that plume. It seems likely that the backscatter continued to dissipate in the 10 minutes between the map taken at 1026 UT and the satellite pass at 1036 UT. These results suggest that TEC depletions might be more enduring than 1-m FAI.

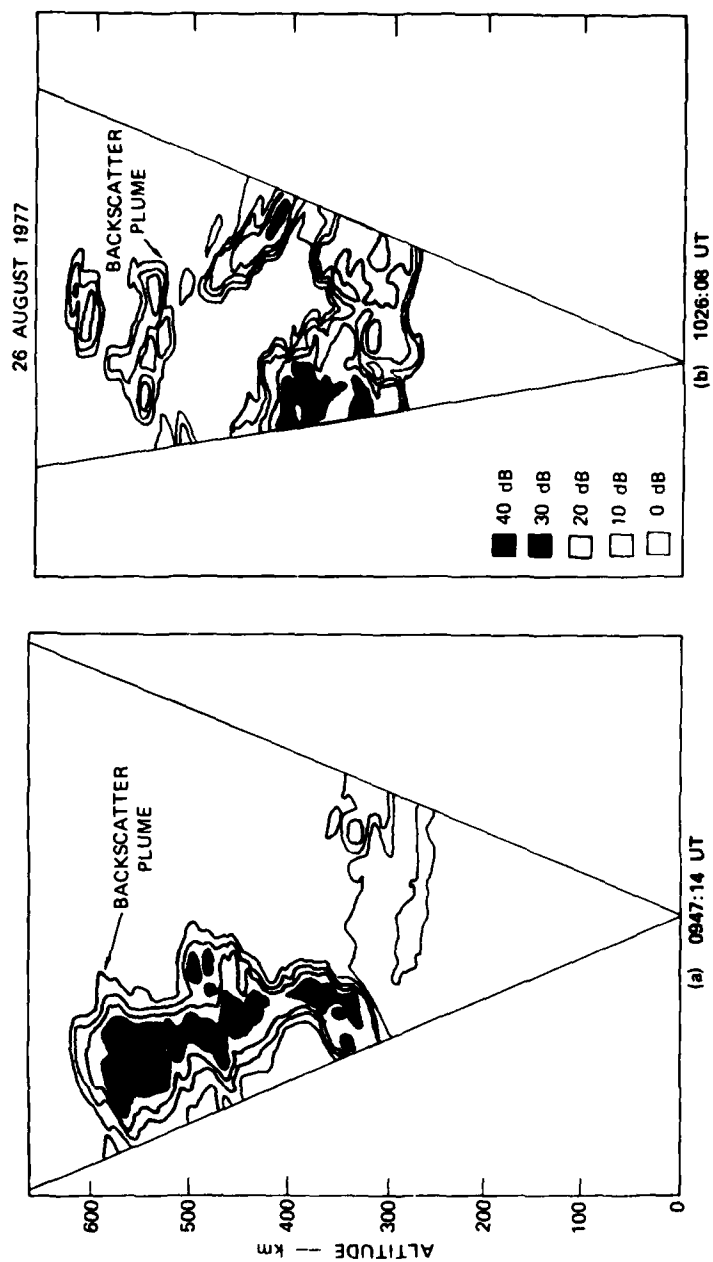


FIGURE 3 TIME SEQUENCE SHOWING THE DECAY OF ALTAIR BACKSCATTER PLUME

IV DISCUSSION AND CONCLUSIONS

The results presented here indicate that backscatter plumes observed with VHF radars are longitudinally coincident with large TEC depletions. Because we know (1) that the altitude of the bottomside F layer was at 250 km, and (2) that the satellite passed at an altitude of 550 km, the TEC depletions must be associated with plasma depletions that occurred within that 300-km altitude interval. Therefore, the TEC depletions must be associated with plasma depletions in the nighttime F layer. On the basis of recent observations of plasma bubbles in the F layer that can be depleted up to three orders of magnitude (Hanson and Sanatani, 1973; McClure et al., 1977), we suggest that the TEC depletions are probably due to plasma bubbles.

Let us consider the implications of the data in terms of a simple plasma-bubble model. The TEC biteout (1034:40 UT) just west of ALTAIR was 35% depleted. That percentage depletion in TEC could also be interpreted as an identical depletion in electron density that extended uniformly from the bottom of the F layer (250 km) to the satellite altitude (550 km). If we assume that the backscatter plume which extends from the bottom of the F layer to an altitude of 450 km defines the volume of the plasma bubble, the electron density within the bubble must be depleted by 53%. Similarly, the TEC depletion of 45% just east of ALTAIR (1035:15 UT) would correspond to a plasma bubble depletion of 69%. In this case, the electron density within the plasma bubble must be $2-3 \times 10^5 \text{ el/cm}^3$.

Another possibility is that the plasma bubble is associated with only the upper portion of a backscatter plume and that the remaining backscatter region is associated with the wake left by the rising plasma bubble (Kelley and Ott, 1978). However, it seems unlikely that the volume of the plasma bubble would be much less than, say, a third of the backscatter region. Since the electron density within the plasma bubble would vary inversely as the volume of the bubble, there are clearly

enough electrons within the plasma bubble to produce radar backscatter. On this basis, it is not unreasonable to argue that radar plumes might be spatially coincident with plasma bubbles. In contrast, it is difficult to envision significant backscatter to occur, for example, from within a plasma bubble that is depleted by three orders of magnitude.

The existence of plasma bubbles without significant electron-density fluctuations (McClure et al., 1977) is also consistent with our finding that the lifetime of radar plumes, or equivalently small-scale FAI, is probably shorter than that of a TEC depletion. The rapid development and decay of 1-m FAI are also consistent with results of comparisons between radiowave scintillations and radar plumes (Morse et al., 1977; Basu et al., 1977; 1978). In particular, Basu et al. (1978) have shown that scintillation-producing (large scale) irregularities coexist with backscatter-producing (small scale) irregularities during the generation phase of equatorial FAI in the evening hours. They further showed that the small-scale FAI decay approximately one hour after sunset but that the large-scale FAI persisted.

The presence of 1-m FAI throughout the entire volume of the plasma bubble is not an accepted relationship by the research community because of two reasons: First, all theoretical production mechanisms for small-scale FAI proposed thus far depend on an electron-density gradient, presumably that associated with the walls of the plasma bubble. For example, the Rayleigh-Taylor instability can be considered to be a gravity-driven analog of the gradient-drift instability, the mechanism believed to produce striations in barium ion clouds (Ossakow and Chaturvedi, 1978). In this mechanism, a rising plasma bubble develops gradients on its top side. Radar backscatter should then be associated with the top wall of plasma bubbles. Other researchers have proposed drift waves as a possible source of small-scale FAI (Hudson and Kennel, 1975; Costa and Kelley, 1978; Huba et al., 1978). Drift waves are also dependent on existing electron-density gradients.

Second, there seems to exist a general misconception that all plasma bubbles are depleted by three orders of magnitude. If one assumes that plasma bubbles are all 99.9% depleted, it is of course difficult

to expect significant backscatter from those regions. It should be stressed that 99.9% depletions reported by Hanson and Sanatani (1973) and McClure et al. (1977) represent the most extreme cases. In our data set, the depletion within the plasma bubble (assumed to be the same volume as the radar plume) was no more than 70%. Such a depletion corresponds to an electron density of $2-3 \times 10^5 \text{ el/cm}^3$ within the depleted region. The only way to increase the percentage depletion within the bubble would be to decrease the volume of the bubble. However, if backscatter is presumed to be associated with electron-density gradients, the size of the backscatter region would be smaller than the size of the plasma bubble.

The above discussion leads us to the conclusion that either (1) the radar plume observed with ALTAIR was spatially coincident with a plasma bubble and that the percentage depletion was about 70%, or (2) the percentage depletion was much larger and that part of the radar plume was associated with regions outside the actual plasma bubble. A possible source of 1-m FAI outside the plasma bubble might be the wake region behind the rising plasma bubble (Kelley and Ott, 1978). Vortices produced in the wake will lead to the development of electron-density irregularities if the eddy motion acts on a background electron-density gradient (the bottom-side and topside of the F layer).

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